

Short-term plant legacy alters the resistance and resilience of soil microbial communities exposed to heat disturbance in a Mediterranean calcareous soil

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1	Title: Short-term plant legacy alters the resistance and resilience of soil microbial									
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24 Abstract

25 Plant legacy is a concept representing the effects exerted by plants on soil once they are no longer growing. We hypothesized that plant species and mixture (intercropping) would induce 26 27 different short-term legacy effects impacting carbon and nitrogen-related soil microbial 28 activities and resistance and resilience after a heat disturbance. A microcosm experiment was 29 conducted using a calcareous Mediterranean soil conditioned by a complete vegetative cycle 30 in a greenhouse with four planting modalities (W = monoculture of Wheat (Triticum aestivum 31 L.); L = monoculture of white Lupin (Lupinus albus L.); WL = both species intercropped; U = 32 unplanted soil). Half of microcosms were incubated at 28° C (C = control conditions) whereas 33 the remaining half were exposed at 48° C for 2 days (S = stress conditions), with an 34 immediately return to control conditions. Microcosms were destructively sampled at 2, 7, 16 35 and 28 days (T2, T7, T16, T28) after the end of the heat disturbance and the following soil 36 measurements were performed: Basal Respiration (BR), Substrate-Induced Respiration (SIR), 37 Nitrification Enzyme Activity (NEA) and N mineral concentrations. Our results demonstrated 38 that monocultures and intercropping promoted different legacy effects under control 39 conditions especially for SIR. WL soils presented lower values of SIR than L and higher than 40 W soils. For SIR, W and WL soils conferred greater resistance to the heat stress, whereas L 41 and WL soils conferred higher resilience at T28. For NEA, no differences between soils were 42 observed for resistance to heat stress, but at T16, soils having WL legacy were more resilient 43 than L soils, but comparable to those having W legacy. Our results highlight that a short-term 44 legacy effect is measurable but greatly differs between C- and N-related microbial activities. 45 We estimated that intercropping had modified ability of soil microorganisms to face heat stress, suggesting that plant legacy effect has to be considered to mitigate extreme climatic 46 47 events in Mediterranean soils.

- **Keywords:** climate change; soil respiration; nitrification; crop mixture; plant soil feedback,
- 49 heat stress.

51 Substantial legacy effects on soil properties are induced by plants and can persist after 52 the disappearance of those plants that were responsible for these effects (van Der Putten et al., 53 2013). Plant legacy effects have been pointed out to influence carbon storage in soils (Lange 54 et al., 2015), succession in plant communities (van der Putten et al., 2016) or plant invasions 55 (Bailey and Schweitzer, 2016). Legacy effects of plants on soils are largely driven by shifts in 56 soil microbial communities. Indeed, plant roots can directly alter soil microbial communities 57 by releasing carbon compounds named rhizodeposits (Eisenhauer et al., 2017; Hunter et al., 2014; Kuzyakov and Blagodatskaya, 2015). Rhizodeposits include a wide range of organic 58 59 compounds including specific root exudates that shape microbial communities in root 60 vicinity, *i.e.* the rhizosphere microbiome (Dennis et al., 2010; Li et al., 2018). Furthermore, 61 composition of rhizodeposits is known to be plant specific (Hunter et al., 2014; Li et al., 62 2018), which suggests that a single species will affect differently soil microbial communities 63 compared to a plant mixture (Chen et al., 2019; Eisenhauer et al., 2017; Tang et al., 2014; 64 Tang et al., 2016). Plants are also able to alter soil microbial communities indirectly by modifying the abiotic properties of the soil, *i.e.* pH, nutrient availability, moisture or soil 65 66 structure (Hinsinger et al., 2009). Several plant legacies were characterized in grasslands 67 (Grman and Suding, 2010; Lange et al., 2015; Strecker et al., 2016; Zak et al., 2003), but 68 debated in agroecosystems while recent publication pointed out its implication in shaping soil 69 microbiome (Li et al., 2018). In agroecosystems, the frequency of perturbations due to annual 70 practices and the low plant diversity barely support comparison to natural ecosystems. 71 However, some agronomic practices such as intercropping are in rupture with conventional 72 agriculture, as it consists to grow two or more species or genotypes in the same field during a significant period of their life cycle (Brooker et al., 2015). Intercropping is of great interest in 73 74 many countries as a solution to replace synthetic inputs by ecological processes (Brooker et

al., 2015; Doré et al., 2011; Malézieux et al., 2009). Long-term legacy effects of intercropping 75 76 have been estimated by Cong et al. (2015), but within 1-year timespan authors are reluctant to 77 expect a significant effect on soil communities. However, Pivato et al. (2017) estimated that 78 after only 12 weeks of bi-species culture, N cycling-related microbial communities were 79 strongly affected especially at low N level compared to each species in monoculture. Grman 80 and Suding (2010) also estimated a short-term legacy effect pointing out the multiple drivers 81 defining both plant legacy duration and intensity. Given the crucial role of soil microbial 82 activities in ecosystem functioning, these plant legacy effects can contribute to agroecosystem 83 stability in the context of climate change and particularly during extreme events. Several 84 studies have been focused on resistance (capacity to withstand change) and resilience 85 (capacity for recovering of function) of C-related soil microbial activities to environmental disturbances (Bérard et al., 2012; Bérard et al., 2011; Chaer et al., 2009; Guillot et al., 2019; 86 87 Hamdi et al., 2011; Mooshammer et al., 2017; Orwin and Wardle, 2005). However, studies 88 about stability of more specialized soil microbial communities like nitrifiers are scarce 89 (Mooshammer et al., 2017; Thion and Prosser, 2014; Wertz et al., 2007). Yet there are 90 reasons to argue that plants are able to shape soil related N-cycle communities (Zak et al., 91 2003). Our rationale is to consider that diverse plant species yield different legacies, 92 ultimately contributing to alleviate effects of extreme climatic events as heat waves in 93 agroecosystems. We hypothesized that intercropped species promote a (i) different legacy 94 effect on soil microbial processes in comparison to single species, and accordingly (ii) 95 influence the resistance and resilience of soils. We focused on nitrification and respiration 96 microbial processes (potential nitrification, substrate-induced respiration and basal 97 respiration) as key processes of the N and C cycles in environmental and plant productivity 98 issues. To achieve this goal, we generated a novel plant legacy in a calcareous cambisol, 99 through a greenhouse experiment conducted with a soil collected in a Mediterranean site near

Montpellier, France. Soil was collected 3 months after a pea crop, sieved at 1-cm to be homogenized and ridded of any coarse organic material, then air dried and stored at ambient temperature before use. In 2017, soil was mixed with perlite and conditioned by one complete vegetative cycle of two crop species grown as single crops (monocultures), (i) white lupin (*Lupinus albus* L.) and (ii) bread wheat (*Triticum aestivum* L.), as well as the mixture of both species (intercrop). Finally, we calculated the resistance and resilience indices (*sensu* Orwin and Wardle (2004)) after a heat disturbance.

107

108 **2. Material and methods**

109 **2.1.** Soil origin and greenhouse experiment before heat stress experiment

110 Soil originated from a plot of the INRA experimental station localized at Mauguio 111 (3°59'6"E, 43°37'13"N, 12 m altitude) in French littoral Mediterranean region. The soil was a 112 calcareous cambisol with a clay loam texture according to USDA classification (clay 294 g kg⁻¹, fine silt 200 g kg⁻¹, coarse silt 233 g kg⁻¹, fine sand 156 g kg⁻¹ and coarse sand 118 g kg⁻¹ 113 114 ¹), which had developed on an alluvial calcareous parent material, typical of the South East 115 plains of Montpellier. The regional climate is typically Mediterranean with surface regularly exposed to drying-rewetting cycles due to irregular rain events mainly concentrated in 116 117 Autumn and Spring. The annual average rainfall at this site during the growing season, from 118 September to June, over the last 20 years was 598 mm. Recorded average soil temperatures at 10 cm depth since 2012 demonstrated that plot had never experienced temperatures higher 119 120 than 38.4 °C. During the four months before soil sampling, 3 cycles of drought were recorded 121 (*i.e.* 20 days with cumulative rainfall lower than 2mm, according to Tang et al. (2014)), while maximal air temperatures ranged from 22.8 °C (in September) to 34.7 °C (in August). The last 122 123 precipitation episode (4.8 mm) preceding soil sampling occurred 5 days earlier. Soil was 124 collected (0-15 cm depth) 3 months after pea harvest, air dried and sieved at 1-cm to homogenize and remove coarse organic material, then finally stored at ambient temperature
before use to preserve inherent soil quality (Salomé et al., 2014).

127 A pot experiment was then conducted in a greenhouse (Centre Mondial de 128 l'Innovation – Roullier Group – Saint-Malo, France) from September to December 2017 with 129 the soil previously described. Before being conditioned in the pots (4-dm³), the soil was re-130 sieved with a 4-mm mesh to remove any organic material and mixed with perlite to allow better conditions for plant root development. Pots were filled with substrate composed of 2 kg 131 of soil and 0.2 kg of perlite. The characteristics of the substrate were clay 202 g kg⁻¹, silt 454 132 g kg⁻¹, sand 343 g kg⁻¹, total CaCO₃ 39 g kg⁻¹, pH_w 8.3, CEC_{Metson} 15 cmol₊ kg⁻¹, organic C 133 10.4 g kg⁻¹, total N 0.95 g kg⁻¹, ammonium (N-NH₄⁺) 20 mg kg⁻¹, nitrate (N-NO₃⁻) 39 mg kg⁻¹. 134 135 The WHC of the substrate corresponded to 28% of gravimetric humidity. For simplicity, substrate will be named herein as soil. The planting modalities of this experiment were: (i) 136 137 monoculture of wheat (W) (Triticum aestivum L. cv. Lennox) at a density of 4 plants per pot, 138 (ii) monoculture of white lupin (L) (Lupinus albus L. cv. Feodora) at a density of 2 plants per 139 pot, (iii) intercropping of both species (WL) at half of the density for each corresponding 140 monoculture (2 wheat and one lupin plants), and (iv) an unplanted soil (U). Each planted and 141 unplanted treatments was replicated 4 times. At maturity, plants (shoots and roots) were 142 harvested and the soil was carefully separated from the roots by gently shaking followed by hand sorting and then kept at 4°C in the dark until microcosm experiment (16 weeks). 143

144

145 **2.2. Heat stress microcosm experiment**

A total of 128 microcosms were prepared (4 replicates x 4 planting treatments x 2 temperatures x 4 sampling dates) containing 50 g of equivalent dry soil from greenhouse experiment, and placed in 1-dm³ glass jars hermetically sealed by a rubber gasket. Soils of microcosms were rehydrated at 100% of the water holding capacity and pre incubated at 12°C

150 for 5 days in order to ensure uniform initial experimental conditions. Afterwards, half of the 151 microcosms (64 samples) were incubated at 28°C and the remaining half at 48°C, 152 corresponding to control (C) and heat stress (S) conditions, respectively. All microcosms were 153 placed into darkness in laboratory incubators (KBWF 720, Binder, Germany) with a precision 154 of 0.4 °C. In order to i) avoid any drying-rewetting cycles during the experiment altering the 155 C and N cycles including nitrification (Fierer and Schimel, 2002) and ii) ensure precise 156 estimation of nitrification potential assays conducted with soil slurries (see section 2.3), we 157 kept soil moisture at 100% WHC. To ensure no variations of water regime in microcosms, a 20-cm³ polypropylene container filled with water was placed in each sealed microcosm. 158 159 Microcosms were regularly weighted to check any soil moisture changes. After two days of 160 heat disturbance at 48°C, the stressed microcosms were transferred back at 28°C. Each 161 microcosm contained an 1M NaOH trap for basal respiration determination (see section 2.3). 162 At each sampling date (T2, T7, T16 and T28 corresponding to 2, 7, 16 and 28 days after the 163 end of the heat stress), samples were immediately split into 4 subsamples for all subsequent 164 soil analyses (substrate-induced respiration (SIR), nitrification enzyme activity (NEA), soil NH4⁺ and NO3⁻ measurements, and soil water content). Measurements of SIR and NEA were 165 166 performed on fresh soils within 2 days after sampling, whereas soil samples for mineral N 167 measurements were frozen (-20°C) until analysis.

168

169 2.3. Soil microbial and chemical analyses

Basal respiration (BR) traps were collected and replaced at each sampling date. An aliquot of 1-cm³ of NaOH was back-titrated with 0.1 M HCl, after adding 2.5-cm³ of BaCl₂ solution (30 %) to precipitate the Na₂CO₃ issued from soil CO₂ respiration. Results of BR were expressed in μ g C-CO₂ day⁻¹ g⁻¹ of dry soil. SIR measurements were performed on 10 g of equivalent dry soil according to Patra et al. (2005), in presence of glucose (1.2 mg C-

glucose g^{-1} dry soil). A CO₂ accumulation kinetic was measured using a gas chromatograph 175 (µGC R3000, SRA instruments, France) at 28°C. SIR was expressed in µg C-CO₂ h⁻¹ g⁻¹ of 176 dry soil. NEA was measured with 3 g of equivalent dry soil according to Patra et al. (2005). 177 Briefly, a $(NH_4)_2SO_4$ solution (50 µg N-NH₄⁺ g⁻¹ dry soil) was added to each soil sample. 178 Samples were placed at 28°C, shaken at 140 rpm to ensure aerobic conditions during 10 179 180 hours, and $NO_2^- + NO_3^-$ production was measured every two hours using a photometer (Smartchem 200, AMS Alliance, France). NEA was expressed in µg of N-(NO₂⁻+NO₃⁻) h⁻¹ g⁻¹ 181 182 of dry soil. Finally, soil ammonium (N-NH4⁺) and soil nitrate (N-NO3⁻) were extracted from a 183 soil:solution ratio of 1:2 with 1M KCl solution after a 30 min of contact at 140 rpm, followed 184 by 30 min of rest. Supernatant were filtered at 0.45 µm and mineral N concentrations were 185 estimated with an Analytical Discrete Multi-Chemistry Analyzer (AQ2, SEAL Analytical, UK). Soil mineral N content were expressed as mg $N-NH_4^+$ and mg $N-NO_3^-$ kg⁻¹ of dry soil. 186

187

188 **2.4. Percentage change relative to control**

In order to investigate recovery rate, percentage change of values in stressed samples relative to those in the controls, at each sampling date, were calculated as follows (Bérard et al., 2011; Chaer et al., 2009):

192

193 relative change (%) =
$$\left[\left(\frac{S_i}{C_i}\right) - 1\right] \times 100$$
 (1)

194

where S_i and C_i are the variable values in the stressed and control soil samples, respectively,
at sampling date i (*i.e.* T2, T7, T16 or T28).

197

198 2.5. Soil microbial resistance and resilience indices

To investigate only potential differences among soils with specific plant legacy (W, L,
WL), unplanted (U) soil was not considered for calculation of resistance and resilience indices
(see section 3.3). Resistance (RS) and resilience (RL) of soil microbial activities to a heat
stress were assessed by the following indices (Orwin and Wardle, 2004):

203

204 RS =
$$1 - \frac{2 \times |D_2|}{C_2 + |D_2|}$$
 (2)

205

where D_2 is the difference between the values of the response variable (BR, SIR or NEA) in control (C₂) and stressed conditions (S₂), 2 days after the end of the disturbance;

208

209 RL =
$$\frac{2 \times |D_2|}{|D_2| + |Dx|} - 1$$
 (3)

210

where D_x is the difference between the control and disturbed soil on 7, 16 or 28 days after the end of the disturbance.

213 Both indices are bounded by -1 and +1, where an index of 1 indicates full resistance 214 (no effect of the disturbance on the response variable) or full resilience (same values of the 215 response variable in control and stressed conditions). For resistance, an index value of 0 216 indicates either 100% decrease or increase in the response variable compared to control, while 217 a negative index represents a modification greater than 100% in the response variable 218 compared to the control (*i.e.* the soil has low resistance). In regard to resilience, an index of 0 219 indicates that the disturbed soil has either not recovered after the end of the disturbance (*i.e.* 220 $D_x = D_2$) or that the stressed soil is different to the control by the same amount, but in the 221 opposite direction (*i.e.* $D_x = -D_2$). A resilience index between 0 and 1 indicates that the response variable has not fully recovered, whereas a negative index value indicates that the absolute value of D_x is higher than the absolute value of D_2 (*i.e.* the soil has low resilience).

224

225 **2.6. Statistical analyses**

226 The effects of planting legacy (*i.e.* both planted (W, L, WL) and unplanted (U) 227 modalities) and heat stress on all soil variables were tested for each sampling date by two-way 228 ANOVA (planting legacy x heat stress). Afterwards, one-way ANOVA was performed in 229 absence of interaction to test differences among treatments. The effect of specific plant legacy 230 (*i.e.* only planted modalities – W, L, WL) on resistance and resilience indices, as well as the 231 effect of planting legacy on percentages of change in microbial activities relative to controls 232 were also tested by one-way ANOVA, within each sampling date. Significant differences 233 between means were tested by Tukey's multiple comparison tests (p < 0.05). Normality of 234 residuals and homogeneity of variance were tested by the Shapiro and Levene's tests, 235 respectively. When necessary, data for microbial activities were squared root-transformed. 236 Principal component analysis (PCA) were performed with all chemical and microbial soil 237 variables. Within a PCA, coordinates of the individuals on each axis were used for testing 238 differences between groups (i.e. levels of a factor) by one-way ANOVA or Kruskal-Wallis 239 tests in case of non-normality. Tukey's multiple comparison tests (p < 0.05) or Wilcoxon rank 240 sum tests (p < 0.05) were then performed accordingly. Finally, a redundancy analysis (RDA) 241 was performed to assess the effect of all soil microbial and chemical variables (SIR, NEA, 242 BR, N-NO₃⁻ and N-NH₄⁺ contents) on resistance (at T2) and resilience (at T16) indices. 243 Statistical analyses were performed with R software v. 3.5.1 (R Core Team, 2018). Packages 244 used for the ANOVA, Tukey's test, PCA and RDA analysis were: car, agricolae, 245 FactoMineR and vegan, respectively.

246

3.1. Contrasted effects of planting and heat stress on microbial activities and soil mineral N content

250 The first two axes of the PCA performed on full data set (BR, SIR, NEA, N-NH₄⁺, N-251 NO₃⁻ measurements for all sampling dates and stress conditions) explained 74.2% of the total 252 variance (Fig. 1). BR, N-NH₄⁺, NEA and N-NO₃⁻ are the major variables contributing to the 253 first axis while SIR was strongly and negatively correlated to the second axis (Table A.1). 254 Both planting legacy (W, L, WL, U) and sampling dates combined with control (C) and stress 255 (S) conditions had a significant impact on values of samples (Kruskal-Wallis p < 0.001; Table 256 A.2 and A.3). Along axes 1 and 2, unplanted soil was significantly different from those with 257 W, L and WL legacies (Wilcoxon test, $p \le 0.001$), with centroids showing negative and 258 positive values, respectively (Fig. 1a). Considering sampling dates combined with control and 259 stress conditions (Fig. 1b), earlier dates were significantly distinct between control and heat 260 stress. However, at these earlier dates, the unplanted stressed soils were closer to their 261 controls, suggesting a greater resistance to heat stress. Interestingly, during the period of 262 resilience, the trajectories of the control and stressed soils evolved along the first axis, and at 263 the later dates (T16, T28) stressed and control soils were not significantly different from each 264 other. At all dates combined, a clear difference was observed between stress and control 265 conditions along the second axis (Wilcoxon test, $p \le 0.001$), with all centroids of the stressed 266 samples showing positive values, whereas centroids of controls showed negative values.

267

3.2. Legacy effect of planting on soil C- and N- related activities in response to heat
stress

Planting legacy was significant for BR values at the first two dates of measure (p <270 271 0.001; Table 1 and Fig C.1). At these dates, L soils clearly differed from unplanted soils, with 272 L soils showing values 1.9-fold greater, when combining control and stress conditions. The 273 planting legacy was still significant at T16, but no effect was observed at T28. An important 274 effect of heat stress was also observed on BR values, particularly at T2 and T7 ($p \le 0.01$ and p275 < 0.001, respectively). At the two first dates, stressed soils presented values 1.35-fold greater 276 than control soils on average. At T16 the difference between stressed and control soils was 277 lower, whereas no significant difference between stressed and control soils was observed at 278 T28.

279 In regard to SIR, planting legacy (p < 0.001), heat stress (p < 0.001) and planting x 280 stress interaction (p < 0.01 at T28 and p < 0.001 at other dates) had a strong effect on SIR 281 values (Table 1 and Fig. C.2). Significant differences between planting legacies of control 282 soils were observed at all sampling dates. However, heat stress alleviated these differences. 283 For instance, at T2 (control), SIR means ranked as follows: L > WL > W = U, whereas no differences were noticed after stress. One should note that L soils exhibited the greatest SIR 284 285 values, while unplanted soils presented the lowest. Conversely to BR, stressed conditions 286 exhibited a dramatic decrease of SIR activities compared to the controls, whatever the 287 sampling date. The most important difference was observed at the first date, when stressed 288 soils were 3.6-fold lower than controls. The L soils were particularly altered at the first date 289 because SIR activity were 5-fold lower than controls. This difference between stress and 290 control conditions was also present at the other dates, but decreased with time.

A very strong effect of planting legacy on NEA values was observed, whatever the sampling date (p < 0.001; Table 1 and Fig. C.3). Unplanted soils exhibited the greatest values compared to planted soils (2.2- to 3.8-fold greater than W, L and WL). The heat stress effect on NEA was particularly observed at T2 and T7 (p < 0.001). At these dates, a significant 295 decrease in NEA values was observed between controls and stressed soils, but to a lesser 296 extent than that observed for SIR. At T16, heat stress had no effect on NEA, whereas at T28, 297 a significant effect of heat stress (p < 0.05) and interaction of heat stress and planting legacy 298 (p < 0.001) were observed. An interesting dynamic was observed from T16 to T28, as NEA 299 values increased particularly for stressed (S) soils, and to a lesser extent for control (C) soils. 300 Indeed, at T28, there was a final ranking with L values becoming closer to U values (U_C > 301 U_S = L_S \geq L_C = W_S = WL_S > WL_C = W_C).

302

303 **3.3. Resistance and resilience of soil microbial activities**

304 Recovery rates are illustrated in Fig. 2. Just after heat disturbance (T2), SIR activity 305 was severely altered (-70%, on average) compared to NEA (-46%) and BR (+36%). The most 306 important decrease at T2 was observed for L soils, with a reduction of 79% in SIR relative to 307 control samples. At this date, L soil also exhibited the largest increase in BR (+59%) relative 308 to control samples. Conversely, for NEA, T2 showed the lowest decrease for the unplanted 309 soil (-15%). Focusing on the other dates for SIR, significant differences between percentages 310 of change were only observed at T7, as follows: U > W = WL > L (Table B.1). For NEA in T16, percentages of change for WL and L, but not for W soils, switched from negative to 311 312 positive. Significant differences between values were observed for NEA at T28. The W soils 313 exhibited greater NEA values in stressed soils relative to controls and compared to WL and L 314 (Table B.1).

Indices of resistance (RS) and resilience (RL) are presented in Table 2. Significant differences in resistance between soils were only observed for SIR (p < 0.05). The W soils exhibited the highest RS index, which was significantly different from L, but not from WL. For resilience indices, significant differences between soils for NEA at T16 (p < 0.01) and for SIR at T28 (p < 0.05) were observed. The RL indices for NEA of both WL and W soils were greater than that of L soils. The RL indices of L soils were negative at T7, T16, and T28,
meaning a very low resilience of NEA. For SIR at T28, the highest RL indices were found for
L and WL soils.

323 The ordination plots of RDA performed on resistance (RS), at T2, and resilience (RL), 324 only at T16, indices clearly discriminated the unplanted soil (Fig. 3 a,b). The resistance 325 indices of NEA were positively associated to NEA and N-NO₃⁻ concentrations in soils, and 326 negatively to N-NH4⁺ contents. The RS indices of SIR were negatively associated to BR. 327 Considering resilience indices of NEA, these latter were negatively linked to NEA and N-328 NO₃⁻, whereas BR resilience indices were negatively linked to SIR and RL of SIR. Finally, 329 considering soils with only specific plant legacies (W, L, WL), L soils slightly distinguished 330 from the other soils on RDA plots for resistance, and strongly differed from the other soils for 331 resilience. Thus, these results confirmed percentages of change (Fig. 2) and RS and RL 332 indices (Table 2). The L soils were clearly different, being characterized by lower RS of SIR 333 and RL of NEA (Fig. 3 a,b).

334

335 **4. Discussion**

4.1. Legacy effect of plants on soil microbial activities under control conditions

337 Marked differences were observed between unplanted and planted soils, suggesting 338 that plants were responsible for a legacy effect. Unplanted soils always exhibited greater 339 values of NEA than those for W, L and WL. As plants are able to shape N-related microbial 340 communities (Moreau et al., 2015; Pivato et al., 2017), it occurred that all planted treatments 341 had a negative legacy effect on the activity of nitrifiers. This negative legacy effect may be 342 attributed to the competition between plants and microorganisms for N (Cantarel et al., 2015; 343 Kuzyakov and Xu, 2013; Xu et al., 2011). According to Schimel and Bennett (2004) plants 344 are effective competitors for ammonium which is also the substrate for nitrifiers, reducing 345 their abundance and activity. Some plant morphological and/or physiological characteristics 346 can also affect the diversity of nitrifiers (Cantarel et al., 2015; Legay et al., 2014; Pommier et 347 al., 2018), or inhibit nitrification (Subbarao et al., 2012). According to these recent results, 348 one should hypothesize that nitrifiers community is shaped by the environment (Erguder et 349 al., 2009; Prosser and Nicol, 2012; Zhalnina et al., 2012), or by plant N-use strategy (Thion et 350 al., 2016). Thereby, we can reasonably suggest a plant legacy effect on nitrification. A great 351 contrast between unplanted and planted soils was also observed for BR and SIR values, with 352 lower values of U compared to W, L and WL. This is consistent with Orwin and Wardle 353 (2005), who observed lower values of soil BR and SIR for bare soil compared to the planted 354 grass treatments. As large amounts of C-compounds can be released by roots as rhizodeposits, 355 e.g. sugars, organic acids, amino acids, mucilage, root or border cells (Kuzyakov and 356 Blagodatskaya, 2015; Philippot et al., 2013), an increase in microbial biomass and activities 357 (Tang et al., 2014) is expected in planted compared to unplanted soils.

358 Considering only planted soils, we also observed different legacy effects among the 359 modalities (W, L, WL), which is consistent with our first hypothesis. Indeed, except for the 360 first sampling date, greater values of NEA were observed for L soil, compared to W and WL, 361 i.e. a stronger legacy effect for lupin than wheat. Le Roux et al. (2013) also reported a 362 positive relationship between the presence of legumes and NEA, which was explained by the 363 build-up of nitrifying communities in response to greater N-NH4⁺ availability in soils, 364 probably due to the release of N derived from Rhizodeposition (NdfR) by legumes (Duchene 365 et al., 2017). In addition, legumes are usually considered less competitive for soil mineral N 366 compared to cereals, which could suggest a higher availability of mineral N in soils under 367 legumes (Mallarino and Wedin, 1990). L soils were also remarkable regarding to C-related 368 microbial variables, showing the greatest values of SIR whatever the date of measurement, as 369 well as the highest values of BR at T2 and T7. White lupin is well known to exude large 370 quantities of organic anions, like citrate (Dissanayaka et al., 2015; Veneklaas et al., 2003; 371 Wang et al., 2010), compared to cereals and other legumes (Wang et al., 2016). Such exudates 372 may have promoted SIR and BR in our study. This positive legacy effect of white lupin on C-373 related microbial variables compared to intercropped wheat and white lupin is however not in 374 agreement with the commonly observed positive effect of plant diversity on soil microbial 375 biomass and respiration (Chen et al., 2019; Eisenhauer et al., 2017; Tang et al., 2014). 376 Accordingly, one should note that legume legacies may be further investigated in 377 agroecosystems in order to decipher their potential role in rotation or in intercropping.

378

379 4.2. Effects of experimental conditions on C-related microbial activities: Basal 380 Respiration and Substrate-Induced Respiration

381 Amongst response variables, SIR was the most affected by heat stress (T2). Samples 382 affected by stress presented reductions varying from 65% to 79% relative to controls. As SIR 383 may be considered as a proxy of the active microbial biomass (Anderson and Domsch, 1978; 384 Orwin and Wardle, 2005), these results suggest a dramatic decrease in microbial biomass by 385 the disturbance. The mortality of temperature-sensitive microbes has been observed in other 386 studies where heat and combined heat-drought stresses were applied (Bérard et al., 2011; 387 Guillot et al., 2019; Hamdi et al., 2011). The effect of temperature on SIR activities was also 388 evident when comparing the heat stress responses of planted soils. In our study, L soils 389 presented the highest values of SIR activity under control conditions, but after heat stress, SIR 390 activity decreased to levels comparable to the other treatments. Then, we concluded that an 391 important microbial biomass in L soils did not confer a better resistance to the heat 392 disturbance. This finding was consistent with Guillot et al. (2019), who observed that 393 Mediterranean soils exhibiting the highest microbial biomass were not systematically 394 associated with a higher resistance. Conversely, W soils which exhibited the lowest values of 395 SIR, presented the greatest resistance index (Table 2), which suggests a greater proportion of 396 more heat-tolerant microbes within this modality. A weaker plant legacy effect on C-related 397 microbial activities due to inherent quality of this type of soil (Salomé et al., 2014) was 398 expected. However, recent studies pointed out that combined heat and drought stresses were 399 able to significantly alter the microbial biomass (Bérard et al., 2011; Guillot et al., 2019), or 400 acclimate of soil communities to heat waves (Bérard et al., 2012) in Mediterranean soils. In 401 our study, the plant-induced effects may be partly due to a soil microbes protection by plant 402 liberation of exopolysaccharides (Kumar and Verma, 2018). Soils with L and WL legacies 403 presented the highest resilience indices and recovery rates for SIR. The rapid progression of L 404 and WL soils over time reduced final differences between these soils and W at T28 (Fig. 2). 405 Despite strongly affected by heat stress, the better resilience of L and WL suggest that heat 406 tolerant microbes rapidly proliferated. The greater resilience of L soils could be attributed to a 407 greater C availability derived from death of the most sensitive microorganisms (Bérard et al., 408 2015), but also by transformation of soil organic matter as induced by the elevated 409 temperature (Bérard et al., 2011; Hamdi et al., 2011). It should be noted that a positive 410 priming effect due to lupin legacy may also contribute to C availability in L soils conferring a 411 greater resilience of L and WL (Wang et al., 2016).

412

413 4.3. Effects of experimental conditions on N-related microbial activities: Nitrification 414 Enzyme Activity (NEA)

415 NEA was affected by heat stress, but to a lesser extent than SIR, while comparable to 416 BR (Fig. 2). Both temperature and moisture are considered as major environmental factors 417 affecting microbial activities and structure. Considerable debate stands in literature about 418 optimal temperature for soil nitrifiers (Taylor et al., 2017) with scarce studies dealing with the 419 effect of nitrification rate at temperatures higher than 40°C. Indigenous nitrifiers have 420 different optimal temperatures depending on native region (Avrahami and Conrad, 2003). 421 Albeit > 40° C temperatures may be experienced in Mediterranean regions, the duration of 422 continuous heat stress, *i.e.* 2 days, is not reproducible under Mediterranean field conditions. 423 French regulation for Mediterranean region defines severe heat waves as a period lasting 2-3 424 days, and characterized by elevated temperature > 35°C and > 21°C during day and night 425 respectively. Thus, night cooling decreases temperature stress even during severe heat waves, 426 which was not the case during our experiment. Then, 2 days on continuous applied 48°C 427 temperature have to be considered as a stress while not fully realistic of field conditions. One 428 should note our moisture content of soils (100% of WHC) during the experiment. Nitrification 429 rate is affected by drying-rewetting cycles (Fierer and Schimel, 2002) such as mineralization 430 (Zhang et al., 2017). Fierer and Schimel (2002) depicted an increase of nitrification potentials 431 in frequently drying-rewetting stress treatments. In their experiment, they do not succeed to 432 attribute any changes of nitrification rates when soil moisture increased from 35% to 50% 433 WHC. Mathieu et al. (2006) estimated no significant differences of nitrification rates using 434 ¹⁵N tracer under 150% WHC (saturated) and 75% of WHC (unsaturated conditions) in soil 435 incubations with 2 mm aggregates. We thus assumed that our uncommon WHC did not alter 436 potential nitrification rates. However, possible restricted oxygen availability due to O_2 437 consumption by mineralization during the incubation period could be pointed out. A O_2 438 restriction may finally alter nitrifier's community growth rate and activity at earlier dates of 439 our experiment. As nitrification potentials are estimated under shaking to provide aeration, 440 oxygen limitation at earlier dates (T2 and T7) may have been less pronounced than at T16. 441 One should note that our goal was to test if planted treatments conferred different legacy 442 effects on potential NEA in response to a heat disturbance. Despite discussed limitations of our experiment, we discussed significant differences while keeping in mind potential 443

limitations of field transposition of our results. We demonstrated that WL legacy exhibited 444 445 comparable resistance and resilience than those with W legacy. The WL soil was the less 446 affected by the heat stress initially (T2), with a reduction by only 31% in NEA relative to the 447 control (Fig. 2), but this trend was not reflected on resistance indices, probably due to the high 448 variability among replicates. However, in terms of resilience, at T16, significant differences 449 between indices were verified, with WL and W soils exhibiting the highest resilience indices. 450 Pivato et al. (2017) estimated a greater abundance of nitrifiers in soils from plant associations, 451 whereas Le Roux et al. (2013) and Zhang et al. (2015) reported a negative effect of plant 452 richness on the abundance of ammonia oxidizers. These latter articles suggested that plant 453 identity rather than plant diversity can affect nitrifier community abundances. Considering 454 gene-based associated groups of nitrifiers namely Ammonia Oxidizing Archaea and Bacteria 455 (AOA and AOB), Thion et al. (2016) estimated greater AOA abundance in the rhizosphere of 456 exploitative grass species, like wheat. Moreover, Taylor et al. (2017) considered that AOA 457 nitrifiers are more tolerant to higher temperatures. Thus, better resilience in planting 458 modalities containing wheat could be attributed to these published results. Finally, one should 459 also note the interesting NEA dynamic observed at later stages in our study. A strong increase 460 in NEA activities of all soils was observed at T28, compared to T16, for both control and heat 461 stress conditions (Table 1). Such increase in NEA at this final date could be explained by a "boost" in the activities of nitrifiers after 30 days of incubation in optimal conditions. Since 462 463 the growth rates of nitrifiers are recognized to be relatively slow (Robertson and Groffman, 464 2015; Wertz et al., 2007), the interpretation of NEA resilience is then highly dependent of the 465 final experimental dates. Accordingly, one should note that individual properties of 466 microorganisms like dormancy, plasticity and growth rate (Shade et al., 2012) are also 467 important components of resistance and resilience of soil microbial communities.

468

469 **Conclusions**

470 Adaptation to climate change of Mediterranean soils regularly exposed to heat 471 waves may be possible through plant legacy effects. Intercropping confers more stability of 472 soil microbial activities compared to single crop species when both C and N are investigated 473 together due to individual benefits of each plant type. Finally, additional studies are essential 474 to investigate if, i) prolonged plant presence will allow a more persistent effect on soils, and if 475 ii) longer period of incubation would be able to detect full recovery of soil microbial activities 476 over time. These results can provide useful information for designing cropping systems, in 477 particular crop rotations and/or intercropping to capitalize on different legacy effects of 478 various crops.

479

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490 **References**

Anderson, J. P. E., and Domsch, K. H. (1978). A physiological method for the quantitative
measurement of microbial biomass in soils. *Soil Biology & Biochemistry* 10, 215-221.

- Avrahami, S., and Conrad, R. (2003). Patterns of community change among ammonia
 oxidizers in meadow soils upon long-term incubation at different temperatures. *Appl. Environ. Microbiol.* 69, 6152-6164.
- Bailey, J. K., and Schweitzer, J. A. (2016). The rise of plant-soil feedback in ecology and
 evolution. *Functional Ecology* 30, 1030-1031.
- Bérard, A., Ben Sassi, M., Kaisermann, A., and Renault, P. (2015). Soil microbial community
 responses to heat wave components: drought and high temperature. *Climate Research*66, 243-264.
- 501 Bérard, A., Ben Sassi, M., Renault, P., and Gros, R. (2012). Severe drought-induced
 502 community tolerance to heat wave. An experimental study on soil microbial processes.
 503 *Journal of Soils and Sediments* 12, 513-518.
- Bérard, A., Bouchet, T., Sévenier, G., Pablo, A. L., and Gros, R. (2011). Resilience of soil
 microbial communities impacted by severe drought and high temperature in the
 context of Mediterranean heat waves. *European Journal of Soil Biology* 47, 333-342.
- Brooker, R. W., Bennett, A. E., Cong, W.-F., Daniell, T. J., George, T. S., Hallett, P. D.,
 Hawes, C., Iannetta, P. P. M., Jones, H. G., Karley, A. J., Li, L., McKenzie, B. M.,
 Pakeman, R. J., Paterson, E., Schöb, C., Shen, J., Squire, G., Watson, C. A., Zhang,
 C., Zhang, F., Zhang, J., and White, P. J. (2015). Improving intercropping: a synthesis
 of research in agronomy, plant physiology and ecology. *New Phytologist* 206, 107117.
- Cantarel, A. A. M., Pommier, T., Desclos-Theveniau, M., Diquélou, S., Schloter, M., and
 Poly, F. (2015). Using plant traits to explain plant–microbe relationships involved in
 nitrogen acquisition. *Ecology* 96(3), 788-799.
- 516 Chaer, G., Fernandes, M., Myrold, D., and Bottomley, P. (2009). Comparative resistance and
 517 resilience of soil microbial communities and enzyme activities in adjacent native
 518 forest and agricultural soils. *Microb Ecol* 58, 414-24.
- 519 Chen, C., Chen, H. Y., Chen, X., and Huang, Z. (2019). Meta-analysis shows positive effects
 520 of plant diversity on microbial biomass and respiration. *Nature communications* 10, 1332.
- 522 Cong, W. F., Hoffland, E., Li, L., Six, J., Sun, J. H., Bao, X. G., Zhang, F. S., and Van Der
 523 Werf, W. (2015). Intercropping enhances soil carbon and nitrogen. *Glob Chang Biol*524 21, 1715-26.
- de Vries, F. T., Liiri, M. E., Bjørnlund, L., Setälä, H. M., Christensen, S., and Bardgett, R. D.
 (2012). Legacy effects of drought on plant growth and the soil food web. *Oecologia* **170**, 821-833.
- Dennis, P. G., Miller, A. J., and Hirsch, P. R. (2010). Are root exudates more important than
 other sources of rhizodeposits in structuring rhizosphere bacterial communities?
 FEMS microbiology ecology 72, 313-327.
- 531 Dissanayaka, D. M. S., Maruyama, H., Masuda, G., and Wasaki, J. (2015). Interspecific
 532 facilitation of P acquisition in intercropping of maize with white lupin in two
 533 contrasting soils as influenced by different rates and forms of P supply. *Plant and Soil*534 **390**, 223-236.
- Doré, T., Makowski, D., Malézieux, E., Munier-Jolain, N., Tchamitchian, M., and Tittonell,
 P. (2011). Facing up to the paradigm of ecological intensification in agronomy:

- revisiting methods, concepts and knowledge. *European Journal of Agronomy* 34, 197-210.
- Duchene, O., Vian, J.-F., and Celette, F. (2017). Intercropping with legume for agroecological
 cropping systems: Complementarity and facilitation processes and the importance of
 soil microorganisms. A review. *Agriculture, Ecosystems & Environment* 240, 148161.
- Eisenhauer, N., Lanoue, A., Strecker, T., Scheu, S., Steinauer, K., Thakur, M. P., and
 Mommer, L. (2017). Root biomass and exudates link plant diversity with soil bacterial
 and fungal biomass. *Scientific reports* 7, 44641.
- 546 Erguder, T. H., Boon, N., Wittebolle, L., Marzorati, M., and Verstraete, W. (2009).
 547 Environmental factors shaping the ecological niches of ammonia-oxidizing archaea.
 548 *FEMS microbiology reviews* 33, 855-869.
- Fierer, N., and Schimel, J. P. (2002). Effects of drying–rewetting frequency on soil carbon
 and nitrogen transformations. *Soil Biology and Biochemistry* 34, 777-787.
- Griffiths, B. S., and Philippot, L. (2013). Insights into the resistance and resilience of the soil
 microbial community. *FEMS Microbiol Rev* 37, 112-29.
- Grman, E., and Suding, K. N. (2010). Within year soil legacies contribute to strong priority
 effects of exotics on native California grassland communities. *Restoration Ecology* 18, 664-670.
- Guillot, E., Hinsinger, P., Dufour, L., Roy, J., and Bertrand, I. (2019). With or without trees:
 Resistance and resilience of soil microbial communities to drought and heat stress in a
 Mediterranean agroforestry system. *Soil Biology & Biochemistry* 129, 122-135.
- Hamdi, S., Chevallier, T., Ben Aïssa, N., Ben Hammouda, M., Gallali, T., Chotte, J.-L., and
 Bernoux, M. (2011). Short-term temperature dependence of heterotrophic soil
 respiration after one-month of pre-incubation at different temperatures. *Soil Biology and Biochemistry* 43, 1752-1758.
- Hinsinger, P., Bengough, A. G., Vetterlein, D., and Young, I. M. (2009). Rhizosphere:
 biophysics, biogeochemistry and ecological relevance. *Plant and soil* 321, 117-152.
- Hunter, P. J., Teakle, G., and Bending, G. D. (2014). Root traits and microbial community
 interactions in relation to phosphorus availability and acquisition, with particular
 reference to Brassica. *Frontiers in plant science* 5, 27.
- Kumar, A., and Verma, J. P. (2018). Does plant—Microbe interaction confer stress tolerance
 in plants: A review? *Microbiological research* 207, 41-52.
- Kuzyakov, Y., and Blagodatskaya, E. (2015). Microbial hotspots and hot moments in soil:
 concept & review. *Soil Biology and Biochemistry* 83, 184-199.
- Kuzyakov, Y., and Xu, X. (2013). Competition between roots and microorganisms for
 nitrogen: mechanisms and ecological relevance. *New Phytologist* 198, 656-669.
- Lange, M., Eisenhauer, N., Sierra, C. A., Bessler, H., Engels, C., Griffiths, R. I., MelladoVázquez, P. G., Malik, A. A., Roy, J., Scheu, S., Steinbeiss, S., Thomson, B. C.,
 Trumbore, S. E., and Gleixner, G. (2015). Plant diversity increases soil microbial
 activity and soil carbon storage. *Nature Communications* 6.
- Le Roux, X., Schmid, B., Poly, F., Barnard, R. L., Niklaus, P. A., Guillaumaud, N., Habekost,
 M., Oelmann, Y., Philippot, L., Salles, J. F., Schloter, M., Steinbeiss, S., and Weigelt,
 A. (2013). Soil environmental conditions and microbial build-up mediate the effect of

- 581 plant diversity on soil nitrifying and denitrifying enzyme activities in temperate 582 grasslands. *PLoS One* **8**, e61069.
- Legay, N., Baxendale, C., Grigulis, K., Krainer, U., Kastl, E., Schloter, M., Bardgett, R. D.,
 Arnoldi, C., Bahn, M., and Dumont, M. (2014). Contribution of above-and belowground plant traits to the structure and function of grassland soil microbial
 communities. *Annals of botany* 114, 1011-1021.
- 587 Li, X., Jousset, A., de Boer, W., Carrión, V. J., Zhang, T., Wang, X., and Kuramae, E. E.
 588 (2018). Legacy of land use history determines reprogramming of plant physiology by
 589 soil microbiome. *The ISME journal*, 1.
- Malézieux, E., Crozat, Y., Dupraz, C., Laurans, M., Makowski, D., Ozier-Lafontaine, H.,
 Rapidel, B., Tourdonnet, S., and Valantin-Morison, M. (2009). Mixing plant species in
 cropping systems: concepts, tools and models. A review. Agronomy for Sustainable
 Development 29, 43-62.
- Mallarino, A., and Wedin, W. (1990). Effect of species and proportion of legume on herbage
 yield and nitrogen concentration of legume grass mixtures. *Grass and Forage Science* 45, 393-402.
- Mathieu, O., Henault, C., Leveque, J., Baujard, E., Milloux, M. J., and Andreux, F. (2006).
 Quantifying the contribution of nitrification and denitrification to the nitrous oxide
 flux using 15N tracers. *Environ Pollut* 144, 933-40.
- Mooshammer, M., Hofhansl, F., Frank, A. H., Wanek, W., Hämmerle, I., Leitner, S.,
 Schnecker, J., Wild, B., Watzka, M., and Keiblinger, K. M. (2017). Decoupling of
 microbial carbon, nitrogen, and phosphorus cycling in response to extreme
 temperature events. *Science advances* 3, e1602781.
- Moreau, D., Pivato, B., Bru, D., Busset, H., Deau, F., Faivre, C., Matejicek, A., Strbik, F.,
 Philippot, L., and Mougel, C. (2015). Plant traits related to nitrogen uptake influence
 plant-microbe competition. *Ecology* 96(8), 2300-2310.
- Orwin, K. H., and Wardle, D. A. (2004). New indices for quantifying the resistance and
 resilience of soil biota to exogenous disturbances. *Soil Biology and Biochemistry* 36,
 1907-1912.
- Orwin, K. H., and Wardle, D. A. (2005). Plant Species Composition Effects on Belowground
 Properties and the Resistance and Resilience of the Soil Microflora to a Drying
 Disturbance. *Plant and Soil* 278, 205-221.
- Patra, A., Abbadie, L., Clays-Josserand, A., Degrange, V., Grayston, S., Loiseau, P., Louault,
 F., Mahmood, S., Nazaret, S., and Philippot, L. (2005). Effects of grazing on microbial
 functional groups involved in soil N dynamics. *Ecological Monographs* 75, 65-80.
- Philippot, L., Raaijmakers, J. M., Lemanceau, P., and van der Putten, W. H. (2013). Going
 back to the roots: the microbial ecology of the rhizosphere. *Nat Rev Microbiol* 11,
 789-99.
- Pivato, B., Bru, D., Busset, H., Deau, F., Matejicek, A., Philippot, L., and Moreau, D. (2017).
 Positive effects of plant association on rhizosphere microbial communities depend on plant species involved and soil nitrogen level. *Soil Biology and Biochemistry* 114, 1-4.
- Pommier, T., Cantarel, A. A., Grigulis, K., Lavorel, S., Legay, N., Baxendale, C., Bardgett, R.
 D., Bahn, M., Poly, F., and Clément, J. C. (2018). The added value of including key

- 624 microbial traits to determine nitrogen related ecosystem services in managed 625 grasslands. *Journal of applied ecology* **55**, 49-58.
- Prosser, J. I., and Nicol, G. W. (2012). Archaeal and bacterial ammonia-oxidisers in soil: the
 quest for niche specialisation and differentiation. *Trends Microbiol* 20, 523-31.
- R Core Team (2018). R: A Language and Environment for Statistical Computing. R
 Foundation for Statistical Computing, Vienna, Austria.
- Robertson, G. P., and Groffman, P. M. (2015). Nitrogen Transformations. *In* "Soil
 microbiology, ecology and biochemistry" (E. A. Paul, ed.), pp. 421-446. Academic
 Press, Burlington, Massachussetts, USA.
- Salomé, C., Coll, P., Lardo, E., Villenave, C., Blanchart, E., Hinsinger, P., Marsden, C., and
 Le Cadre, E. (2014). Relevance of use-invariant soil properties to assess soil quality of
 vulnerable ecosystems: The case of Mediterranean vineyards. *Ecological Indicators*43, 83-93.
- 637 Schimel, J. P., and Bennett, J. (2004). Nitrogen mineralization: challenges of a changing
 638 paradigm. *Ecology* 85, 591-602.
- Shade, A., Peter, H., Allison, S. D., Baho, D. L., Berga, M., Burgmann, H., Huber, D. H.,
 Langenheder, S., Lennon, J. T., Martiny, J. B., Matulich, K. L., Schmidt, T. M., and
 Handelsman, J. (2012). Fundamentals of microbial community resistance and
 resilience. *Front Microbiol* 3, 417.
- 643 Strecker, T., Macé, O. G., Scheu, S., and Eisenhauer, N. (2016). Functional composition of
 644 plant communities determines the spatial and temporal stability of soil microbial
 645 properties in a long-term plant diversity experiment. *Oikos* 125, 1743-1754.
- Subbarao, G., Sahrawat, K., Nakahara, K., Ishikawa, T., Kishii, M., Rao, I., Hash, C., George,
 T., Rao, P. S., and Nardi, P. (2012). Biological nitrification inhibition—a novel
 strategy to regulate nitrification in agricultural systems. *In* "Advances in agronomy",
 Vol. 114, pp. 249-302. Elsevier.
- Tang, X., Bernard, L., Brauman, A., Daufresne, T., Deleporte, P., Desclaux, D., Souche, G.,
 Placella, S. A., and Hinsinger, P. (2014). Increase in microbial biomass and
 phosphorus availability in the rhizosphere of intercropped cereal and legumes under
 field conditions. *Soil Biology and Biochemistry* **75**, 86-93.
- Tang, X., Placella, S. A., Daydé, F., Bernard, L., Robin, A., Journet, E.-P., Justes, E., and
 Hinsinger, P. (2016). Phosphorus availability and microbial community in the
 rhizosphere of intercropped cereal and legume along a P-fertilizer gradient. *Plant and Soil* 407, 119-134.
- Taylor, A. E., Giguere, A. T., Zoebelein, C. M., Myrold, D. D., and Bottomley, P. J. (2017).
 Modeling of soil nitrification responses to temperature reveals thermodynamic
 differences between ammonia-oxidizing activity of archaea and bacteria. *ISME J* 11,
 896-908.
- Thion, C., and Prosser, J. I. (2014). Differential response of nonadapted ammonia-oxidising
 archaea and bacteria to drying-rewetting stress. *FEMS Microbiol Ecol* 90, 380-9.
- Thion, C. E., Poirel, J. D., Cornulier, T., De Vries, F. T., Bardgett, R. D., and Prosser, J. I.
 (2016). Plant nitrogen-use strategy as a driver of rhizosphere archaeal and bacterial
 ammonia oxidiser abundance. *FEMS Microbiol Ecol* 92.

- van der Putten, W. H., Bardgett, R. D., Bever, J. D., Bezemer, T. M., Casper, B. B., Fukami,
 T., Kardol, P., Klironomos, J. N., Kulmatiski, A., Schweitzer, J. A., Suding, K. N.,
 Van de Voorde, T. F. J., Wardle, D. A., and Hutchings, M. (2013). Plant-soil
 feedbacks: the past, the present and future challenges. *Journal of Ecology* 101, 265276.
- van der Putten, W. H., Bradford, M. A., Pernilla Brinkman, E., van de Voorde, T. F., and
 Veen, G. (2016). Where, when and how plant-soil feedback matters in a changing
 world. *Functional Ecology* 30, 1109-1121.
- Veneklaas, E. J., Stevens, J., Cawthray, G. R., Turner, S., Grigg, A. M., and Lambers, H.
 (2003). Chickpea and white lupin rhizosphere carboxylates vary with soil properties and enhance phosphorus uptake. *Plant and Soil* 248, 187-197.
- Wang, B. L., Tang, X. Y., Cheng, L. Y., Zhang, A. Z., Zhang, W. H., Zhang, F. S., Liu, J. Q.,
 Cao, Y., Allan, D. L., Vance, C. P., and Shen, J. B. (2010). Nitric oxide is involved in
 phosphorus deficiency-induced cluster-root development and citrate exudation in
 white lupin. *New Phytologist* 187, 1112-1123.
- Wang, X., Tang, C., Severi, J., Butterly, C. R., and Baldock, J. A. (2016). Rhizosphere
 priming effect on soil organic carbon decomposition under plant species differing in
 soil acidification and root exudation. *New Phytologist* 211, 864-873.
- Wertz, S., Degrange, V., Prosser, J. I., Poly, F., Commeaux, C., Guillaumaud, N., and Le
 Roux, X. (2007). Decline of soil microbial diversity does not influence the resistance
 and resilience of key soil microbial functional groups following a model disturbance. *Environ Microbiol* 9, 2211-9.
- Ku, X., Ouyang, H., Richter, A., Wanek, W., Cao, G., and Kuzyakov, Y. (2011). Spatio temporal variations determine plant–microbe competition for inorganic nitrogen in an
 alpine meadow. *Journal of Ecology* 99, 563-571.
- Zak, D. R., Holmes, W. E., White, D. C., Peacock, A. D., and Tilman, D. (2003). Plant
 diversity, soil microbial communities, and ecosystem function: are there any links?
 Ecology 84, 2042-2050.
- Chalnina, K., de Quadros, P. D., Camargo, F. A., and Triplett, E. W. (2012). Drivers of
 archaeal ammonia-oxidizing communities in soil. *Front Microbiol* 3, 210.
- Chang, N. N., Sun, Y. M., Wang, E. T., Yang, J. S., Yuan, H. L., and Scow, K. M. (2015).
 Effects of intercropping and Rhizobial inoculation on the ammonia-oxidizing
 microorganisms in rhizospheres of maize and faba bean plants. *Applied Soil Ecology*85, 76-85.
- Zhang, W., Liang, C., Kao-Kniffin, J., He, H., Xie, H., and Zhang, X. (2017). Effects of
 drying and wetting cycles on the transformations of extraneous inorganic N to soil
 microbial residues. *Scientific reports* 7, 9477.

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705 **Figure and table captions**

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Fig. 1. Principal component analysis (PCA) performed on basal respiration (BR), substrateinduced respiration (SIR), nitrification enzyme activity (NEA) and soil mineral N content as nitrate (N-NO₃⁻) and ammonium (N-NH₄⁺). Symbols stand for planting legacy (W – wheat, L - lupin, WL – wheat + lupin, U – unplanted) (a) and sampling dates (T2, T7, T16 and T28 – 2, 7, 16 and 28 days after the end of the heat stress) within control (C) and heat stress (S) conditions (b). Centroids are represented by larger symbols.

Fig. 2. Percentages of change in microbial activities (BR – basal respiration, SIR – substrateinduced respiration, NEA – nitrification enzyme activity) relative to controls along all sampling dates after the end of the heat stress (T2, T7, T16, T28). Symbols represent planting legacy (W – wheat, L – white lupin, WL – wheat + lupin, U – unplanted). Vertical bars indicate standard errors (n=4). Asterisks indicate significant differences between planting legacies within a sampling date (one-way ANOVA, followed by a Tukey post hoc test (p <0.05)).

Fig. 3. Redundancy analysis (RDA) based on resistance (RS) (a) or resilience (RL) (b) indices of basal respiration (BR), substrate-induced respiration (SIR) and nitrification enzyme activity (NEA), constrained by BR, SIR, NEA, as well as soil mineral N content as nitrate (N-NO₃⁻) and ammonium (N-NH₄⁺). Resistance indices were calculated for T2, whereas resilience indices were calculated for T16 (2 and 16 days after the end of the heat stress). Groups of planting legacies are represented by W (wheat) in blue, L (lupin) in orange, WL (wheat + lupin) in green and U (unplanted) in red. The RDA was performed on center-reduced matrix.

727 **Table 1.** Basal respiration (BR – in μ g C-CO₂ day⁻¹ g⁻¹ dry soil), substrate-induced respiration (SIR – in μ g C-CO₂ h⁻¹ g⁻¹ dry soil) and nitrification enzyme activity (NEA – in μ g N-NO₂⁻ + 728 729 N-NO₃⁻ h^{-1} g⁻¹ dry soil) of each planting legacy (W – wheat, L – lupin, WL – wheat+lupin, U 730 - unplanted), within control and stress conditions, for each sampling date after the end of the 731 heat stress (T2, T7, T16, T28). Values are the means of four replicates (n=4). Thresholds of 732 probability of the two-way ANOVA (Planting x Stress) are indicated at the top of each 733 variable-related data set. When significant interaction between factors (Planting x Stress) is 734 present, different bold letters represent significant differences between 8 means (i.e. control 735 and heat stress conditions combined; Tukey post hoc test (p < 0.05)). Different lowercase 736 letters represent significant differences within a single column, whereas different uppercase 737 letters represent either significant differences between stress conditions (Control vs Stress), or 738 between planting legacies with control and stress conditions combined (Tukey post hoc test (p 739 < 0.05)).

740 Table 2. Indices of resistance and resilience based on basal respiration (BR), substrate-741 induced respiration (SIR) and nitrification enzyme activity (NEA) for each specific plant 742 legacy (W – wheat, L – lupin, WL – wheat+lupin). Resistance indices were calculated on T2 743 date (2 days after the end of the heat stress), whereas resilience indices were calculated on T7, 744 T16 and T28 dates of measure (7, 16 and 28 days after the end of the heat stress). Values are 745 the means of four replicates (n=4). Probabilities of the one-way ANOVA are indicated at the 746 bottom of each variable-related data set. Different letters represent significant differences 747 within a column (Tukey post hoc test (p < 0.05)).

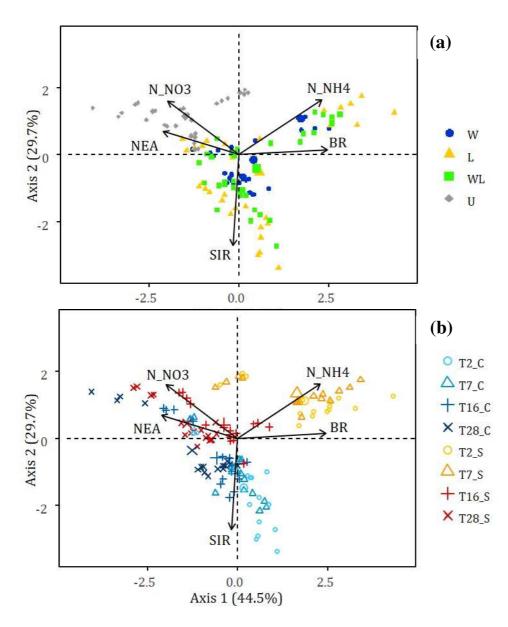


Fig. 1. Principal component analysis (PCA) performed on basal respiration (BR), substrate-induced respiration (SIR), nitrification enzyme activity (NEA) and soil mineral N content as nitrate $(N-NO_3)$ and ammonium $(N-NH_4)$. Symbols stand for planting legacy (W – wheat, L – lupin, WL – wheat + lupin, U – unplanted) (a) and sampling dates (T2, T7, T16 and T28 – 2, 7, 16 and 28 days after the end of the heat stress) within control (C) and heat stress (S) conditions (b). Centroids are represented by larger symbols.

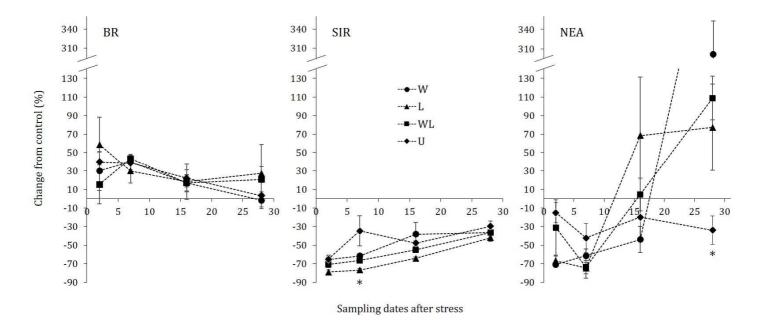


Fig. 2. Percentages of change in microbial activities (BR – basal respiration, SIR – substrate-induced respiration, NEA – nitrification enzyme activity) relative to controls along all sampling dates after the end of the heat stress (T2, T7, T16, T28). Symbols represent planting legacy (W – wheat, L – white lupin, WL – wheat + lupin, U – unplanted). Vertical bars indicate standard errors (n=4). Asterisks indicate significant differences between planting legacies within a sampling date (one-way ANOVA, followed by a Tukey post hoc test (p < 0.05)).

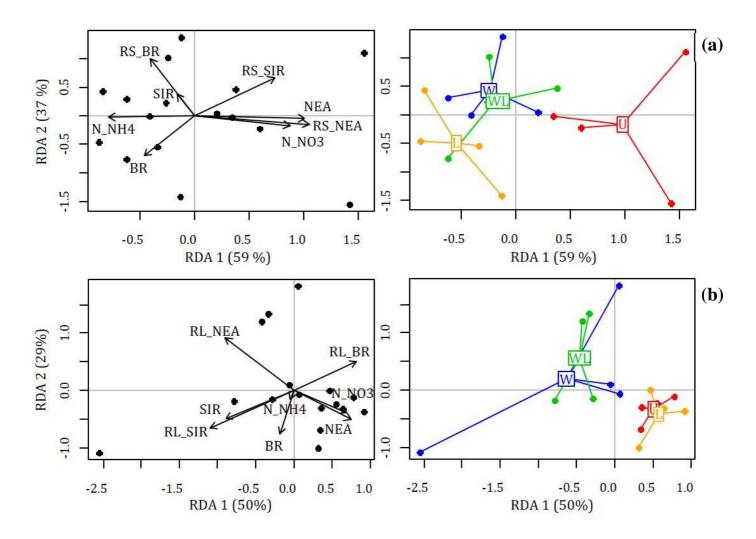


Fig. 3. Redundancy analysis (RDA) based on resistance (RS) (a) or resilience (RL) (b) indices of basal respiration (BR), substrate-induced respiration (SIR) and nitrification enzyme activity (NEA), constrained by BR, SIR, NEA, as well as soil mineral N content as nitrate (N-NO₃⁻) and ammonium (N-NH₄⁺). Resistance indices were calculated for T2, whereas resilience indices were calculated for T16 (2 and 16 days after the end of the heat stress). Groups of planting legacies are represented by W (wheat) in blue, L (lupin) in orange, WL (wheat + lupin) in green and U (unplanted) in red. The RDA was performed on center-reduced matrix.

		Т	2		Т	7		T	16		T2	28
Variable	Factor											
	Planting legacy	< 0.001***		< 0.001***		< 0.05*		ns				
	Heat stress	< 0.01** ns		< 0.001*** ns		< 0.05* ns		ns ns				
	Planting x Stress											
DD		Control ^B	Stress ^A		Control ^B	Stress ^A	_	Control ^B	Stress ^A	_	Control	Stress
BR	W	32.82 ab	41.39 ab	BC	21.52 b	29.79 bc	В	20.18 a	22.73 a	AB	20.07 a	19.31 a
	L	41.61 a	63.79 a	A	37.91 a	48.14 a	А	21.79 a	24.76 a	А	20.19 a	23.88 a
	WL	39.32 a	43.41 ab	AB	21.56 b	38.46 ab	В	20.99 a	24.37 a	А	15.78 a	18.37 a
	U	23.73 b	33.23 b	С	18.63 b	25.30 c	В	16.17 a	19.21 a	В	16.50 a	16.12 a
	Planting legacy	< 0.001***		< 0.001***		< 0.001***		< 0.001***				
	Heat stress	< 0.00	< 0.001***		< 0.0	< 0.001***		< 0.001***		< 0.001***		
	Planting x Stress	< 0.001***		< 0.001***		< 0.001***		< 0.01**				
SIR		Control	Stress		Control	Stress	_	Control	Stress	_	Control	Stress
SIK	W	5.77 c	2.01 d		4.33 c	1.65 e		4.21 bc	2.64 de		4.94 bc	3.15 de
	L	12.42 a	2.59 d		9.11 a	2.10 e		7.69 a	2.75 de		6.92 a	4.00 cd
	WL	9.24 b	2.78 d		6.56 b	2.20 de		5.38 b	2.41 de		5.38 b	3.43 de
	U	4.98 c	1.72 d		3.50 cd	1.71 e		3.47 cd	1.81 e		3.82 d	2.67 e
	Planting legacy	< 0.001***		< 0.001***		< 0.001***		< 0.001***				
	Heat stress	< 0.001***		< 0.001***		ns		< 0.05*				
	Planting x Stress	ns		ns		ns		< 0.001***				
		Control ^A	Stress ^B		Control ^A	Stress ^B	_	Control	Stress	_	Control	Stress
NEA	W	0.54 b	0.16 b	В	0.40 b	0.15 b	В	0.59 b	0.30 b	В	0.65 c	2.69 bc
	L	0.54 b	0.17 b	В	0.57 ab	0.12 b	В	0.74 b	1.12 ab	В	2.83 bc	4.66 b
	WL	0.42 b	0.28 b	В	0.40 b	0.11 b	В	0.43 b	0.45 ab	В	1.28 c	2.63 bc
	U	1.44 a	1.21 a	А	0.91 a	0.49 a	А	1.99 a	1.36 a	А	8.63 a	5.38 b

Table 1. Basal respiration (BR – in μ g C-CO₂ day⁻¹ g⁻¹ dry soil), substrate-induced respiration (SIR – in μ g C-CO₂ h⁻¹ g⁻¹ dry soil) and nitrification enzyme activity (NEA – in μ g N-NO₂⁻ + N-NO₃⁻ h⁻¹ g⁻¹ dry soil) of each planting legacy (W – wheat, L – lupin, WL – wheat+lupin, U – unplanted), within control and stress conditions, for each sampling date after the end of the heat stress (T2, T7, T16, T28). Values are the means of four replicates (n=4). Thresholds of probability of the two-way ANOVA (Planting x Stress) are indicated at the top of each variable-related data set. When significant interaction between factors (Planting x Stress) is present, different bold letters represent significant differences between 8 means (i.e. control and heat stress conditions combined; Tukey post hoc test (p < 0.05)). Different lowercase letters represent significant differences within a single column, whereas different uppercase letters represent either significant differences between stress conditions combined (Tukey post hoc test (p < 0.05)).

		T2	Τ7	T16	T28
Variable	Plant legacy	(Resistance)		(Resilience)	
	W	0.52	- 0.00	0.27	0.48
BR	L	0.34	0.31	0.55	0.27
	WL	0.60	- 0.22	0.20	0.26
p-value (or	ne-way ANOVA)	0.61	0.31	0.61	0.86
	W	0.21 a	0.16	0.47	0.35 b
SIR	L	0.11 b	0.16	0.32	0.54 a
	WL	0.17 ab	0.19	0.37	0.53 a
p-value (one-way ANOVA)		0.05*	0.87	0.63	0.04*
	W	0.17	0.21	0.24 a	- 0.63
NEA	L	0.20	-0.03	-0.58 b	- 0.48
	WL	0.29	- 0.10	0.46 a	- 0.68
p-value (or	ne-way ANOVA)	0.34	0.16	0.008**	0.56

Table 2. Indices of resistance and resilience based on basal respiration (BR), substrate-induced respiration (SIR) and nitrification enzyme activity (NEA) for each specific plant legacy (W – wheat, L – lupin, WL – wheat+lupin). Resistance indices were calculated on T2 date (2 days after the end of the heat stress), whereas resilience indices were calculated on T7, T16 and T28 dates of measure (7, 16 and 28 days after the end of the heat stress). Values are the means of four replicates (n=4). Probabilities of the one-way ANOVA are indicated at the bottom of each variable-related data set. Different letters represent significant differences within a column (Tukey post hoc test (p < 0.05)).